



Distinct methanotrophic communities exist in habitats with different soil water contents



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ABSTRACT

Aerobic methane oxidizing bacteria play a key role in controlling global climate by reducing methane (CH₄) emissions in natural ecosystems. We studied the community assembly processes and co-occurrence interactions of soil methanotrophs in three habitats (an alpine meadow, a marsh meadow, and a marsh) from Qinghai-Tibetan Plateau. Methanotrophic communities and CH₄ oxidation potentials varied considerably between the habitats, and the diversity of methanotrophs was significantly lower in marsh meadow than in the other two soils ($P < 0.001$). Methanotrophic bacterial diversity was significantly correlated with soil dissolved organic carbon (DOC), pH, total carbon (TC), and total nitrogen (TN), while methanotrophic community structure was mostly correlated with soil C/N, TC, soil moisture, and TN. Stochasticity dominated methanotrophic community assembly, and increased from 67.6% in the alpine meadow and 68.0% in the marsh meadow to 98.2% in the marsh. The natural connectivity of co-occurrence network was greater in the alpine meadow than in the other two habitats, suggesting a more stable network in the alpine meadow. Methanotroph diversity contributed to the sub-network topological differences and keystone species were identified such as USC_γ, *Methylobacter*, and RPC-1. The results suggest the existence of distinct community assembly processes and co-occurrence patterns of soil methanotrophs among different habitats, which may ultimately enhance the understanding of factors influencing CH₄ oxidation rates.

1. Introduction

Wetlands contribute nearly one-third of the global methane emissions, releasing about 250 Tg of methane per year to the atmosphere (Mitsch et al., 2013), thus serving as a major carbon reservoir with potential to accelerate climate change (IPCC, 2013). Aerobic methane oxidizing bacteria (MOB) are ubiquitous in wetlands (Deng et al., 2013), forest soils (Kolb et al., 2005), and cold environments, such as tundra soils, permafrost sediments, and arctic soils (Trotsenko and Khmelenina, 2005; Liebner and Wagner, 2007; Martineau et al., 2010), and play a critical role in global climate control and element cycling (Hanson and Hanson, 1996; Bodelier and Laanbroek, 2004; Murrell,

2010; Reeceburgh, 2014). They can oxidize one-carbon (C1) compounds, such as methane and methanol using the particulate (pMMO) and/or soluble (sMMO) methane monooxygenase enzymes (Hanson and Hanson, 1996), and are currently restricted to the phylum *Proteobacteria*. Methanotrophs can be divided into two groups. Type I methanotrophs are members of the *Gammaproteobacteria*, and primarily use the ribulose monophosphate cycle for carbon assimilation. Type II methanotrophs are *Alphaproteobacteria*, which use the serine cycle for carbon assimilation (Hanson and Hanson, 1996). The pMMO is present in all known methanotrophs except the genera *Methylocella* and *Methyloferula* (Vorobev et al., 2011), while these two genera seem to play a minor role in habitats on the Tibet Plateau. (Deng et al., 2013). The

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pmoA gene encodes the β -subunit of the pMMO, which exist in most well known methanotrophs, and is an excellent marker in identifying methanotrophs (Mcdonald et al., 2008; Knief, 2015). Previous molecular surveys have revealed dynamic changes in methanotrophic community structure, abundance, and activity. For example, combining real-time quantitative PCR (qPCR) and denaturing gradient gel electrophoresis (DGGE) analysis, Yun et al. (2010, 2012) observed that *Methylobacter* (type I methanotrophs) species outnumbered *Methylocystis* (type II methanotrophs) in the Zoige wetland (at the eastern edge of Tibetan Plateau) and their abundance and activity varied between different soil conditions. Cloning and sequencing analyses from Riganqiao peatlands (belong to Zoige wetland) revealed high type II methanotroph abundance in the area and their diversity was distinct between hummocks and hollows (Deng et al., 2013). The two sites have different soil properties and somewhat different vegetation types. In general, the Riganqiao soil is more acidic and has both higher soil organic carbon and higher total nitrogen than the Zoige soil. In addition to pH, factors such as vegetation type, oxygen, nutrient availability, and soil moisture are also important for the selection of methanotroph species (Hanson and Hanson, 1996). In fact, Type I and type II MOB have been shown to display markedly different behavior, which can be interpreted as ecologically different strategies. Type I methanotrophs (e.g., *Methylobacter*, *Methylosarcina*, *Methylomonas*) would be classified as competitors or ruderals. They can quickly become dominant in various environments. Sufficient nutrient availability seems to be a prerequisite for the proliferation of type I methanotrophs; however, type II methanotrophs are at an advantage in niches where resources are more limiting (Siljanen et al., 2011; Ho et al., 2013). Nevertheless, there has been limited study of methanotroph abundance and community structure across habitats differing in water content.

Soil microbial community structure is shaped by a combination of deterministic and stochastic processes (Chase and Myers, 2011). Deterministic processes are niche-based, including environmental filtering, biotic interactions, and interspecific trade-offs that enable species to co-exist within communities for long periods of time (Stegen et al., 2012). Stochastic processes essentially involve random processes such as reproduction, dispersal, colonization (Barber and Marquis, 2011), and ecological drift (Bell, 2001). Both deterministic and stochastic processes simultaneously govern the spatial distribution of microbial communities, and different processes are dominant in different cases (Caruso et al., 2011; Dumbrell et al., 2010; Peay and Bruns, 2014). Therefore, anthropogenic activities (e.g. global warming caused by greenhouse gases) may influence soil microbial assembly through the modulation of both deterministic and stochastic processes (Zhang et al., 2016). However, to the best of our knowledge, few data are available regarding MOB community assembly.

In natural ecosystems, microorganisms living together form complex networks through positive (e.g. mutualistic) and negative (e.g. competitive) interactions, as well as interactions with no impacts on fitness (Faust and Raes, 2012). Network analysis has been used to explore interactions among microorganisms in various habitats, including soils (Lupatini et al., 2014), lakes (Xue et al., 2018), and the human gut (Faust et al., 2012). The topological properties of the networks reflect the interactions between microorganisms that traditional methods could not have predicted. For example, taxa in the co-occurrence network with high closeness centralization and low betweenness centrality can be considered as keystone taxa (Berry and Widder, 2014). The scores (network topology properties) also change as environmental conditions shift (Wu et al., 2016). Owing to such merits, network analysis has become a powerful tool in the study of microbial ecology.

The Qinghai-Tibetan Plateau (QTP) is the Earth's largest ($2 \times 10^6 \text{ km}^2$) and highest (average $\sim 4500 \text{ m a. s. l.}$) plateau. It has been called the “third pole” of the Earth, and is a major CH_4 emission hotspot with annual emissions estimated to be 0.56–1.00 Tg (Ding and Cai, 2007). We collected soil samples from three habitats (alpine meadow, marsh meadow, and marsh) in a continuum water-content

region. In this study, we hypothesized that the methanotrophic community assembly processes and co-occurrence network patterns would be distinct between the three habitats with different water contents. The objective of this study was to investigate the factors influencing soil methanotrophic community diversity, structure and community assembly processes on a local scale between habitat types.

2. Materials and methods

2.1. Study site description

Three transects were set-up across an alpine meadow, a marsh meadow and a marsh in a single valley in the northeast ($37^{\circ}50'N37^{\circ}60'N$, $101^{\circ}32'E \sim 101^{\circ}35'E$) of QTP (Fig S1A; Table S1). This area ($\sim 10 \text{ km}^2$ square sub-region) represents the world's largest high-altitude (about 3200 m) and low-latitude permafrost area. Mean annual temperature (MAT) in the valley is about -2°C and mean annual precipitation (MAP) is about 500 mm. The area is dominated by meadow vegetation, but also contains alpine meadow, marsh meadow and marsh (Fig S1B). Soil type was mat-cryic cambisol in alpine meadow, mol-cryic cambisol in marsh meadow and organic cryic gleysol in marsh; soil texture was silty clay soil, loam clay soil, and loam clay soil, respectively. As our sampling site, we choose a continuum in geography containing the three vegetation types. The plant community is dominated by *Kobresia humilis* (C. A. Mey. ex Trautv.) Sergiev., *K. pygmaea* C. B. Clarke., and *Ajania tenuifolia* in the alpine meadow; *Stipa capillata* Linn., *Scirpus pumilus* Vahl., and *Saussurea romuleifolia* Franch in the marsh meadow; and *Heleocharis dulcis* (Burm. f.) Trin., *Carex* spp., and *Crepis flexuosa* (Ledeb.) C. B. Clarke in the marsh.

2.2. Sampling and soil physicochemical analysis

Samples were collected on August 26th, 2016. In total, 45 soil samples were collected from the three habitats and three transects. On each transect, 15 samples (5 from each habitat site) were collected from the top 0–10 cm of the soil. The samples were then transported to the laboratory in insulated boxes fitted with sterile ice bags. They were divided into three parts as follows: one part was stored at 4°C for the measurement of soil physicochemical properties; one was air dried and sieved through a 2 mm mesh for the incubation experiment; and one was stored at -40°C for DNA extraction. Soil pH was determined using a pH meter (Thermo Orion-868, Waltham, MA, USA) after shaking a soil water (1:5 w/v) suspension. Soil moisture was determined gravimetrically. The water-filled pore space (% WFPS) was calculated based on the results of gravimetric water content and the soil bulk density values reported by Hu et al. (2008). Total carbon (TC) and total nitrogen (TN) were measured using an elemental analyzer (Vario MAX, Elementar, Germany). Dissolved organic nitrogen (DON) and dissolved organic carbon (DOC) were determined from a mixture with a ratio of 10 g fresh soil to 100 ml 2 M KCl (deionized water for DOC) by shaking for 1 h, after standing for 1 h and then filtered (Fisher G4 1.2 μm). Then the extracted was determined using a continuous flow analyzer (San + System, Skalar, Holland).

2.3. CH_4 oxidation potential

Before incubation, the soil was pre-incubated at 60% water holding capacity (WHC) at 28°C in darkness for four days under ambient air conditions. Aliquots (5 g, wet weight) of the pre-incubated soil were then placed into 120 ml serum vials capped with butyl rubber stoppers and CH_4 was injected into the gas headspace of the vials to yield final mixing ratios of 2% (Siljanen et al., 2011). Triplicate flasks were prepared and a control serum vial without soil was used to confirm the absence of gas leaks. CH_4 concentrations were measured daily with a gas chromatograph (Agilent 7890, Santa Clara, CA, USA) equipped with a flame ionization detector using 1 ml gas samples from the bottle

headspaces. The concentration of CH₄ began to decrease immediately at the beginning of incubation (Fig S2). Potential CH₄ oxidation rates were determined from the slope of the linear regression equation of CH₄ oxidation versus incubation time (Shrestha et al., 2012).

2.4. DNA extraction, *pmoA* gene quantification and 454 pyrosequencing

Fresh soil, 0.5 g, was used for total DNA extraction using a FastDNA[®] Spin kit (MO Bio, Carlsbad, CA, USA) according to the manufacturer's instructions. The DNA was further purified using DNA Clean-Up kits (MO Bio, Carlsbad, CA, USA). Final DNA yield was determined using a Nanodrop 1000 spectrophotometer (Thermo Scientific, Wilmington, DE). The A189F/mb661r primer pair was used to amplify the *pmoA* gene (Knief, 2015). Soil samples were then sequenced on a Roche Genome Sequencer FLX System platform (454 Life Science, Branford, CT, USA). The absolute abundance of *pmoA* genes was determined using qPCR assays, which were performed in 20 µl volumes containing 10 µl Premix Ex Taq (RR420A, SYBR[®] Premix Ex Taq[™], Takara Bio Inc, Japan), 0.2 µl of each primer, and 2 µl of DNA template. All reactions were performed in triplicate on a CFX96 Optical Real-Time Detection System (Bio-Rad, Laboratories Inc., Hercules, CA, USA). Cycling conditions were set as follows: 30 s at 95 °C, followed by 39 cycles of denaturation at 95 °C for 10 s, 55 °C for 30 s, 72 °C for 30 s, and 80 °C for 5 s. The standard was obtained by a 10-fold dilution series of a *pmoA* gene clone from a stock concentration of 2.83×10^9 copies. The amplification efficiency was between 94.8% and 98.1% with R² values of 0.999–1, no signals were detected in the negative controls.

2.5. Bioinformatic analysis

The raw sequences obtained from the present study were deposited into the NCBI Sequence Read Archive (SRA) with the accession number SRP127483.

A total of 274,405 sequences were obtained from 45 soil samples. QIIME 1.9.1 was used to analyze the sequencing data (Caporaso et al., 2010). Chimeric reads were identified and removed using USEARCH 6.1 (Edgar et al., 2011). Frameshifts were corrected using FrameBot (Wang et al., 2013) with the online version of the FunGene Pipeline (Fish et al., 2013). High-quality sequences were clustered into operational taxonomic units (OTUs) with UCLUST algorithm (Edgar, 2010) using the pick *de novo* mode at a 7% distance cutoff value (Degelmann et al., 2010). The *pmoA* gene annotation was performed in standalone BLAST 2.7.1 + according to a database with 6628 sequences (Dumont et al., 2014). Database sequences were aligned using MUSCLE software (Edgar, 2004). A phylogenetic tree was constructed using the FastTree software program (Price et al., 2010).

2.6. Statistical analysis

SPSS Statistics 23.0 (IBM Corporation, Armonk, NY, USA) was used to perform ANOVA analysis to evaluate differences in soil physiochemical properties, CH₄ oxidation potential, and gene copies among habitats. Non-metric multidimensional scaling (NMDS), Simper analysis, Mantel test, Partial Mantel test, and pairwise comparisons (ADONIS, ANOSIM and MRPP) were conducted with the vegan package in R3.4.1 (<http://cran.stat.sfu.ca/>). Weighted Unifrac and unweighted Unifrac distances were used to evaluate phylogenetic dissimilarity and were calculated with QIIME script (*beta_diversity_through_plots.py*) (Caporaso et al., 2010). Partial least squares path models (PLS-PM) were constructed with the function “inner plot” using the plspm package of R3.4.1 (Sanchez, 2013). In addition, we used multiple regressions on distance matrices (MRM) measuring the relative explained variances of soil physiochemical properties, CH₄ oxidation potential, and methanotroph diversity to explain the community structure. MRM analyses were conducted with the ecodist package of R3.4.1.

To determine the assembly processes driving the community

composition within a sample, we calculated the Nearest Taxon Index (NTI) values of each sample and beta Nearest Taxon Index (βNTI) for paired samples. NTI values were calculated in R3.4.1 using the picante package (Webb et al., 2008), and βNTI values were calculated using the comdist functions in phylocom 4.2 (Webb et al., 2008). If the NTI values are positive and mean values are above 0 ($P < 0.05$), co-occurring species are more affected by phylogenetic clustering than by overdispersal; conversely, negative values and mean values below 0 ($P > 0.95$) indicate that co-occurring species are more affected by overdispersal (Webb, 2000). If βNTI values are > 2 or < -2 , deterministic processes are critical in shaping community composition, while if βNTI values are between -2 and $+2$, stochastic processes play a key role (Stegen et al., 2012). To further quantify the stochastic differences, we estimated pairwise community turnover using the Raup-Crick metric extended to incorporate species' relative abundances (RC_{bray}). Pairwise comparisons between communities that did not deviate from the null model distribution were evaluated as the contribution of dispersal limitation ($|\beta\text{NTI}| < 2$ and $\text{RC}_{\text{bray}} > +0.95$) and homogenizing dispersal ($|\beta\text{NTI}| < 2$ and $\text{RC}_{\text{bray}} < -0.95$); RC_{bray} values between -0.95 and $+0.95$ indicate that compositional turnover between a given pair of communities is ‘undominated’ (Stegen et al., 2013).

Network structures were calculated in R3.4.1 using the WGCNA package and visualized using the interactive platform Gephi 0.9.1 using directed network and the Fruchterman-Reingold layout (Bastian et al., 2009). We considered a valid co-occurrence event to have a Spearman's correlation coefficient (ρ) > 0.6 ($P < 0.01$; Junker and Schreiber, 2008). We generated sub-networks for each soil sample from networks by preserving OTUs presented in each site using subgraph functions in igraph packages. The correlation between sub-network topological features and environmental factors were tested using ANOVA analysis. The Z_i-P_i thresholds were based on metabolic network methods (Guimerà and Amaral, 2005). Briefly, we sorted all species into four groups: peripherals ($z_i \leq 2.5$; $p_i \leq 0.62$), connectors ($z_i \leq 2.5$; $p_i > 0.62$), module hubs ($z_i > 2.5$; $p_i \leq 0.62$), and network hubs ($z_i > 2.5$; $p_i > 0.62$) (Olesen et al., 2007).

3. Results

3.1. Soil chemical properties and CH₄ oxidation activity

In all three transects, soil chemical properties were significantly different among the three habitats, and there were almost no differences when comparing the same habitat type among replicate transects (Table S2). As expected, soil moisture was significantly higher in marsh soil ($265 \pm 83\%$) than marsh meadow soil ($211 \pm 48\%$) and alpine meadow soil ($48 \pm 6\%$); these correspond to the soil water-filled pore space (WFPS) contents of 84.91% in marsh, 80.38% in marsh meadow, and 54.23% in alpine meadow soil. Soil DOC, TC, and TN in the marsh meadow were all significantly higher ($P < 0.001$, ANOVA) than those in the alpine meadow and marsh. The soil C/N ratio and DON were higher in the marsh than in the alpine meadow and marsh meadow. Soil pH exhibited no variation across the three habitats. The potential CH₄ oxidation rate was significantly higher in marsh meadow soil ($15.60 \pm 4.04 \mu\text{g g}^{-1}$ dry soil h^{-1}) than the alpine meadow ($1.80 \pm 0.50 \mu\text{g g}^{-1}$ dry soil h^{-1}) and marsh soils ($13.70 \pm 0.26 \mu\text{g g}^{-1}$ dry soil h^{-1}), with similar values among replicate transects when comparing a single habitat (Table S2). The *pmoA* gene copy numbers were significantly higher in the marsh ($5.86 \pm 3.51 \times 10^7 \text{g}^{-1}$ dry soil) than those in the alpine meadow ($1.58 \pm 0.39 \times 10^7 \text{g}^{-1}$ dry soil) and the marsh meadow ($2.39 \pm 1.02 \times 10^7 \text{g}^{-1}$ dry soil). Methanotrophs typically contain two copies of the *pmo* operon in their genome, and therefore *pmoA* abundance can be used as proxy (at a ratio of 2:1) for cell abundance. Based on this calculation, the specific CH₄ oxidation potential in the marsh meadow was significantly higher ($P < 0.05$) than those in the

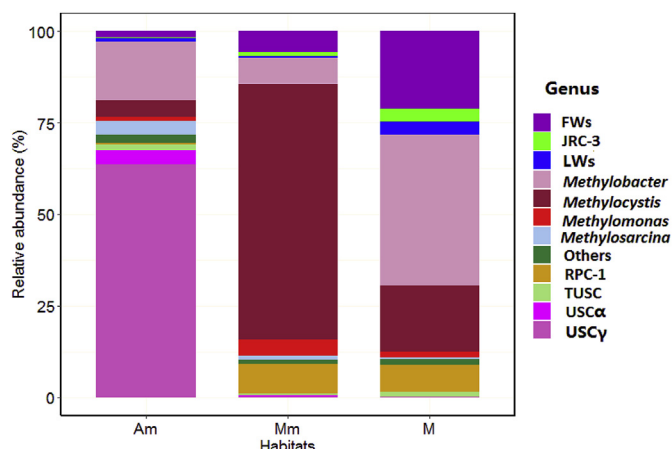


Fig. 1. The relative abundances of dominant soil methanotrophic genera in alpine meadow, marsh meadow, and marsh. Methanotroph taxonomy is based on Lüke and Frenzel (2011). The abbreviations of lineages: USC, upland soil cluster; FWs, freshwater sediment of Lake Wintergreen, Michigan, USA; RPC, rice paddy cluster; LW, sediment of Lake Washington, USA; JRC, Japanese rice cluster; TUSC, tropical upland soil cluster. Am: alpine meadow; Mm: marsh meadow; M: marsh.

other habitats; however, there was no significant difference between alpine meadow and marsh (Fig. S2).

3.2. Variations in methanotrophic community composition among alpine meadow, marsh meadow, and marsh

Soil methanotrophic community composition differed between the alpine meadow, marsh meadow, and marsh (Fig. 1) and there were almost no differences among three replicate transects when comparing the same habitat (Table S3). The methanotroph genera were divided into five groups according to their phylogenetic classification. Generally, type Ib, type IIa, and type Ia methanotrophs dominated the communities and overall amongst the samples they accounted for 69.5%, 69.8%, and 43.9% of the species, respectively (Table S4). The relative abundances of each lineage varied between the habitats. The USC γ (type Ib) were the dominant species in the alpine meadow and accounted for 66.3% of the methanotrophic community, but were rare in the marsh meadow (0.1%) and the marsh (0.02%). The relative abundance of *Methylocystis* (type IIa) was higher in the marsh meadow (69.7%) than in the alpine meadow (4.7%) and the marsh (18.0%). In the marsh, *Methylobacter*, FWs, and *Methylocystis* were dominant species (Fig. 1). RPC-1 (type Ib) were well represented in the marsh meadow (7.9%) and in the marsh (7.6%) compared with in the alpine meadow (0.4%). In contrast, *Methylosarcina* (Type Ia) had higher relative abundance in the alpine meadow (3.7%) than in the marsh meadow (1.2%) and the marsh (0.4%).

3.3. Environmental drivers of CH₄ oxidation potential, soil methanotroph diversity and community structure

CH₄ oxidation potential and *pmoA* gene copies were positively correlated with soil C/N in all soil samples (Table S11) and increased from the drier to the wetter soils (Fig. S3). Community diversity (Shannon index, Phylogenetic diversity, and OTU richness) were lower in the marsh meadow than in the alpine meadow and the marsh (Fig. 2A–C) and there were almost no differences among transects (Table S5). NMDS demonstrated that soil methanotrophic community structure varied among habitats (Fig. 2D), which was further demonstrated by Simper analysis (Table S6). Community dissimilarity was significantly different between each compartment, which was further confirmed by ADONIS, ANOSIM, and MRPP analyses based on Bray-

Curtis distances (Table S7, S8). “Within-site pairs” and “between-site pairs” of Bray-Curtis dissimilarity revealed that community dissimilarity was greater between habitats than between transects (Figs. S4, S5).

Compared with all other factors, soil DOC had the highest negative correlation with soil methanotroph diversity (Table S10). A Mantel test between community similarity and phylogenetic similarity with geographical distance, and environmental dissimilarity revealed that environmental variables played key roles in shaping community similarity and phylogenetic similarity overall, and that geographical distance significantly influenced soil methanotrophic community similarity in the alpine meadow and the marsh (Table S12). After eliminating the effect of geographic distance, partial Mantel test results revealed that environmental factors were significantly correlated with methanotroph community similarities ($r = 0.356$, $P = 0.001$) and phylogenetic similarities (Unweighted Unifrac: $r = 0.447$, $P = 0.001$; Weighted Unifrac: $r = 0.291$, $P = 0.001$) across all soil samples. Environmental variables played a critical role in shaping the community similarity and phylogenetic similarity, particularly in the marsh meadow (Bray-Curtis: $r = 0.357$, $P = 0.006$; Weighted Unifrac: $r = 0.251$, $P = 0.046$) and the marsh (Unweighted Unifrac: $r = 0.473$, $P = 0.014$) but not in the alpine meadow. In addition, when we eliminated the effect of environmental distance, a highly significant correlation between geographic distance and methanotrophic community dissimilarity ($r = 0.266$, $P < 0.001$) was detected across all soil samples, and independently in the alpine meadow ($r = 0.231$, $P < 0.001$) and the marsh ($r = 0.218$, $P = 0.031$) (Table 1). PLS-PM and MRM analyses were carried out to reveal the potential pathways influencing methanotrophic community structure with regard to three factors, including physicochemical properties, CH₄ oxidation potential, and soil methanotroph alpha diversity (Fig. 3A–C). After retaining the most significant variables, the PLS path modeling explained 89% of the community variation. The direct effect of soil parameters (path coefficient = 0.703) on methanotrophic community structure was greater than the direct effects of CH₄ oxidation potential (path coefficient = 0.366) and soil methanotroph alpha diversity (path coefficient = 0.038) (Fig. 3A). Among the environmental variables, soil C/N had the highest explained variance (0.54) (Fig. 3C), suggesting that soil C/N was a key driver explaining methanotrophic community structure (Fig. 5A; Table S9).

3.4. Quantifying influences of ecological processes in shaping methanotrophic community structure

In general, we observed that all the mean NTI values were above zero in habitats and transects (Fig. S6; Table S13), which indicated that methanotrophic communities were phylogenetically clustered. The values of β NTI metrics were between -2 and $+2$ either within or between community comparisons (Fig. S7A, B), which indicated that a stochastic process mainly governed methanotrophic community dynamics, and the relative contribution of stochasticity was lower in alpine meadow (67.6%) and marsh meadow (68.0%), compared to marsh (98.2%) (Fig. 4A and B; Table S14). The stochastic processes were different between the three habitats. Compositional turnover was undominated in alpine meadow (74.3%) and marsh (77.9%); there was a relatively high rate of dispersal limitation in the alpine meadow (25.7%) while no dispersal limitation was detected in marsh meadow and marsh; however, there was a high rate of homogenizing dispersal in marsh (22.1%) and marsh meadow (97.1%) and no homogenizing dispersal was detected in the alpine meadow (Fig. 4C). Moreover, the β NTI metric values in the alpine meadow and the marsh meadow were significantly lower in soils from the marsh ($P < 0.001$, ANOVA), these two habitats have a proportion of deterministic processes (32.4%, 32.0%, respectively). To assess the relative influence of environmental variables on β NTI, Mantel tests between the β NTI metric and the environmental factors were performed. Among the environmental variables, phylogenetic null deviation was mostly correlated to the changes

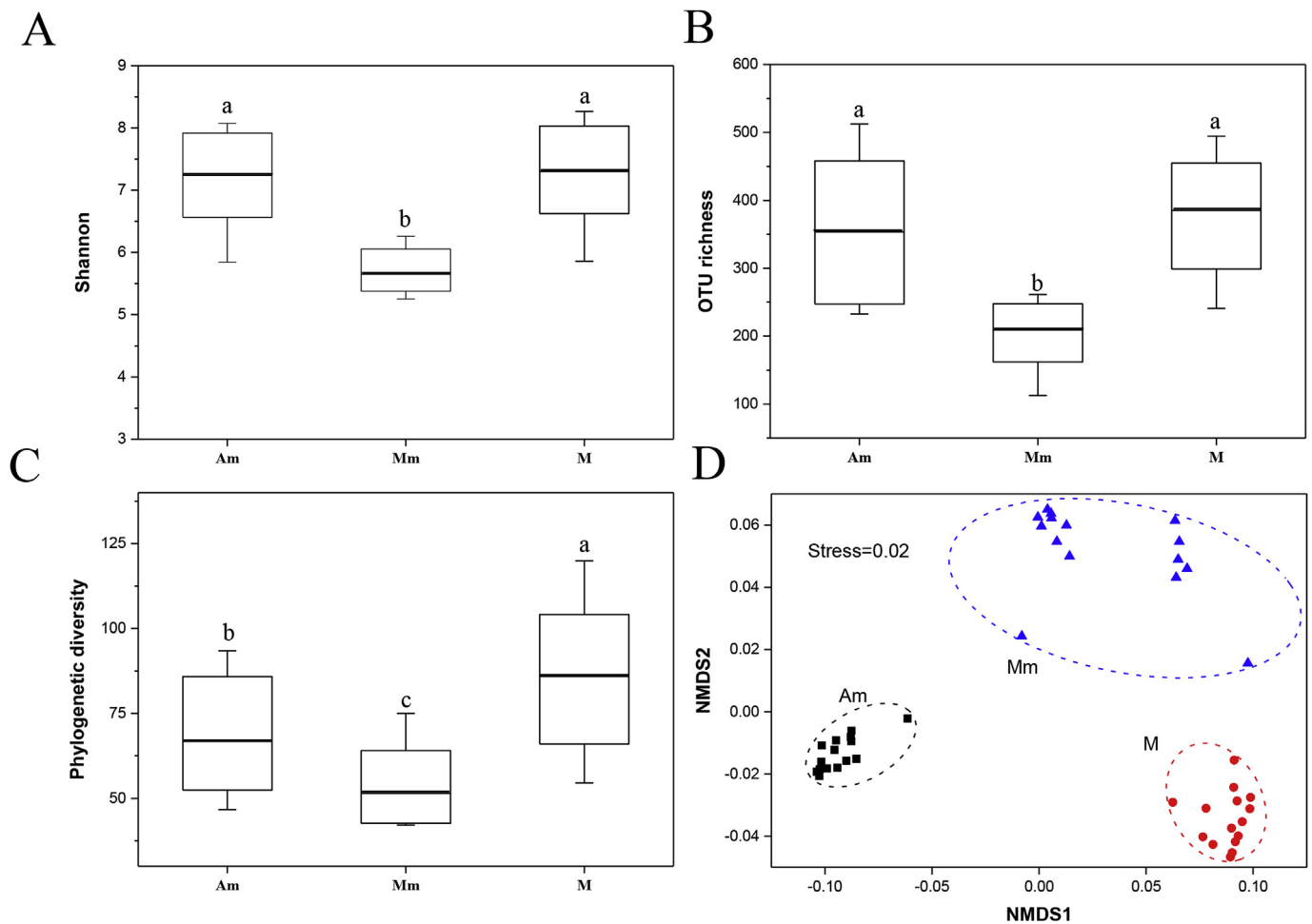


Fig. 2. Shannon index (A), OTU richness (B), and phylogenetic diversity (C) of soil methanotrophic communities in the three habitats. Values in the columns that do not share the same letter differ significantly ($P < 0.05$). Soil methanotrophic community structure in the three habitats as indicated by non-metric multidimensional scaling (NMDS) plots (D). Am: alpine meadow; Mm: marsh meadow; M: marsh.

in DOC (Table S15; Fig. S8; $r = 0.140$, $P = 0.001$). This could suggest that the filtering effect of soil DOC (e.g. increase or decrease in the soil DOC variation between sites) could be causing the disturbances in the phylogenetic structure pattern across sites, leading to the phylogenetic conservation of methanotrophic communities within sites.

3.5. Topological properties and keystone species in methanotroph communities co-occurring in alpine meadow, marsh meadow and marsh

In order to characterize and evaluate the co-occurrence interactions of the methanotrophic communities in each of the habitats, we

constructed co-occurrence networks and calculated the topological properties of the networks and sub-networks (Table S17). Considering the number of correlations, we found that predominantly positive correlations in all habitats, which indicated that methanotrophs exhibited cooperative rather than competitive relationships (Oliveira et al., 2014). Alpine meadow had the highest node connectivity (i.e., average degree of 10.346) and the highest average path length (1.64) compared with marsh meadow and marsh (Table S16). The diameters of the networks displayed a similar trend, peaking in the alpine meadow (7 edges) with the lowest value observed in the marsh (3 edges). Topological patterns of degree distribution showed that *pmoA*

Table 1

Partial Mantel tests were conducted to compare the relative effects of environment and spatial distance on soil methanotrophic community compositional similarity and phylogenetic similarity. Env.dist is the environmental heterogeneity matrix calculated with Euclidean and Geog.dist is the geographic distance matrix. Bold values indicate significant differences ($***P < 0.001$; $**P < 0.01$; $*P < 0.05$). Am: alpine meadow; Mm: marsh meadow; M: marsh.

Habitat	Composition similarity		Phylogenetic similarity			
	Bray-Curtis		Unweighted Unifrac		Weighted Unifrac	
	Effect of Env.dist	Effect of Geog.dist	Effect of Env.dist	Effect of Geog.dist	Effect of Env.dist	Effect of Geog.dist
	Controlling for Geog.dist	Controlling for Env.dist	Controlling for Geog.dist	Controlling for Env.dist	Controlling for Geog.dist	Controlling for Env.dist
Am	-0.029	0.231***	-0.08	0.108	-0.242	0.058
Mm	0.357***	0.254	0.127	0.018	0.251*	0.008
M	0.047	0.218*	0.473*	-0.008	0.148	0.046
All samples	0.356***	0.266***	0.447***	0.171**	0.291***	0.154**

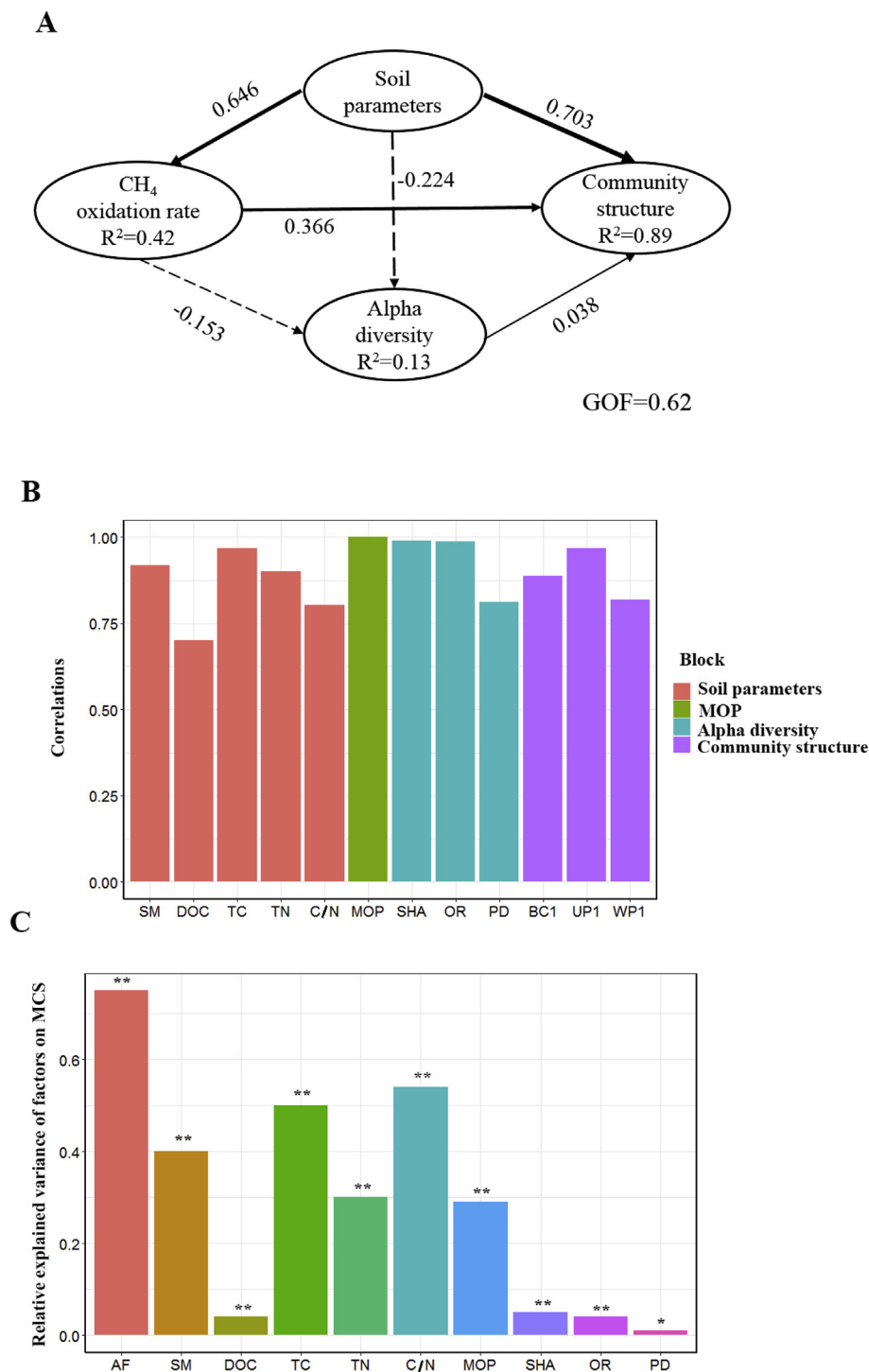


Fig. 3. Path analysis diagrams for soil methanotrophic community structure. The weights of the arrows indicate the strengths of the causal relationships, supplemented by a path coefficient, continuous and dashed arrows indicate positive and negative relationships, respectively (A); the amount of variability explained by the latent variable (B); and the relative explained variance of factors on soil methanotrophic community structure (MCS) in all sampling sites (C). AF: All factors; SM: Soil moisture; DOC: dissolved organic carbon; TC: total carbon content; TN: total nitrogen; C/N: carbon-to-nitrogen ratio; MOP: methane (CH₄) oxidation potential; SHA: Shannon index; OR: OTU richness; PD: PD whole tree; BC1: Bray Curtis PcoA1; UP1: Unweighted unfracPcoA1; WP1: Weighted unfracPcoA1.

genes in all the three habitats exhibited a typical power-law distribution (Fig. S10), which indicated a scale-free network structure and not a random co-occurrence pattern. When comparing network stability among the three habitats, less fluctuation from natural connectivity patterns was observed in response to the proportions of removed nodes in the alpine meadow (Fig. 5B), which indicated a more stable network topology.

We further generated sub-networks for each soil sample by keeping OTUs associated with specific samples and all edges among them in the co-occurrence network in each habitat. A number of subnetwork-level topological features were calculated and separated into three clusters based on hierarchical cluster analyses on the dissimilarities (Fig. S9).

Cluster 1 included node number, edge number, global clustering coefficient, and average clustering coefficient. Cluster 2 included edge density and degree centralization. Cluster 3 included clusters, eigenvector centralization, modularity, and betweenness centralization. After evaluating their correlations with the environmental variables, we observed that the edge numbers, similar to other topological features in the cluster, such as node number, global clustering coefficient, and average clustering coefficient, had a positive correlation with soil methanotroph alpha diversity, with a lower correlation coefficient in the marsh than in the other two habitats. Edge density and degree centralization in cluster 2 were negatively correlated with soil methanotroph alpha diversity in all the three habitats. Additionally, cluster

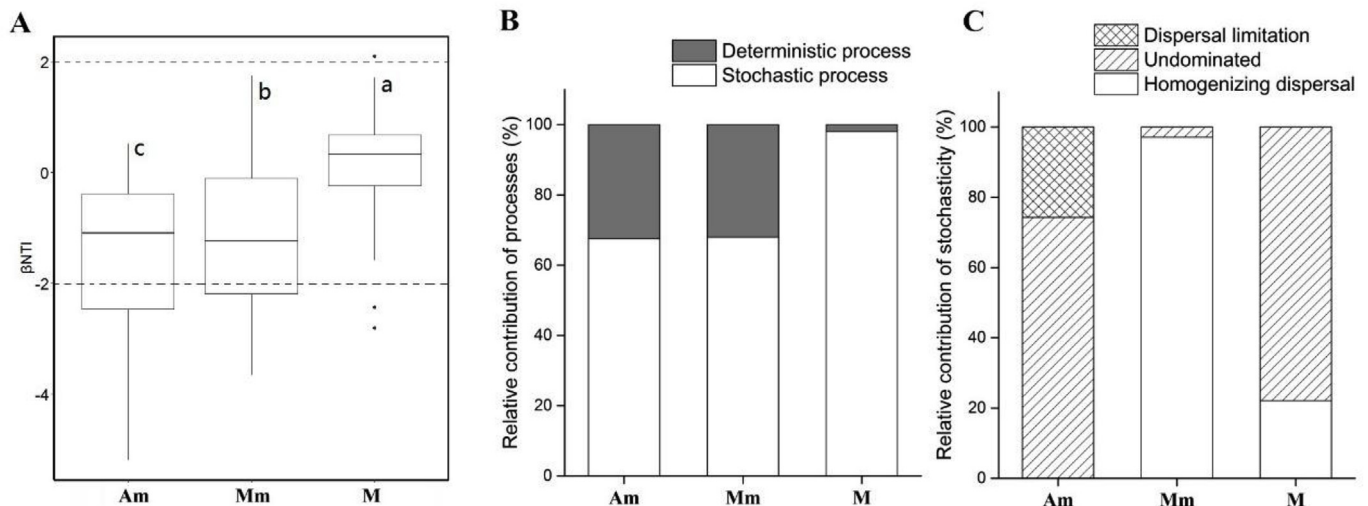


Fig. 4. Methanotrophic community assemblages tested by β NTI values (A). (-2 and $+2$ were the ecological processes threshold for deterministic and stochastic force). Values in the columns that do not share the same letter differ significantly ($P < 0.05$); the relative contribution of each processes (B); and the relative contribution of ecological processes to stochasticity in the three habitats (C). Am: alpine meadow; Mm: marsh meadow; M: marsh.

number, eigenvector centralization, and modularity were positively correlated with soil alpha diversity. Betweenness centralization had a negative relationship with methanotroph alpha diversity in the marsh meadow and the marsh (Table S18).

The Zi-Pi relationships among OTUs revealed that most nodes were peripherals in all habitats, a few nodes were connectors in every habitat (Table S19), while there were no module hubs or network hubs. We identified the ‘connectors’ as the keystone species in the networks. *Methylobacter* (OTU 158) and USC γ (OTU 2013) were connectors in the alpine meadow. There were relatively more low abundance OTUs (Table S20) serving as connectors in the marsh meadow, e.g. OTU 3746, OTU 2909, OTU 962, OTU 6544, OTU 3608 (all assigned to USC γ), and OTU 3154 (assigned to RPC-1) than in the alpine meadow and the marsh meadow. OTU 6431 (assigned to *Methylobacter*) and OTU 3149 (Unclassified methanotrophs) were detected as connectors in the marsh network (Fig. 6).

4. Discussion

4.1. Potential CH_4 oxidation rates and different groups of soil methanotrophs

Soil CH_4 oxidation potential is typically concentrated at the soil-water interface. Therefore, CH_4 oxidation potential was higher in the marsh meadow and the marsh (high water table areas) than in the alpine meadow (Fig. S2), which is consistent with the findings of a previous study (Rachwal et al., 2014). CH_4 oxidation potential was spatially positively correlated with C/N and soil moisture (Table S11), which is also consistent with the results of several studies (Roslev and King, 1994; Martineau et al., 2014). In the present study, soil gravimetric moisture exhibited an increasing trend from the alpine meadow (48%), the marsh meadow (211%) to the marsh (265%), which was correlated with a significant increase in soil CH_4 oxidation activity. Soil structure and soil pore system influence water regimes and gas fluxes. A high-water table makes conditions more anoxic. Consequently, the methanogenic activity was high, which facilitated aerobic methanotrophy near the oxic interface (Smith et al., 2003; Wagner, 2017).

This study demonstrated that the methanotrophic communities vary among the three habitats with distinct soil physiochemical properties optima. Some species in genus *Methylocystis*, a group of type II methanotrophs, have been reported to have a relatively high affinity for CH_4 (Baani and Liesack, 2008; Limbri et al., 2014). In the present study,

a high abundance of *Methylocystis* was detected in the marsh meadow. High abundances of *Methylocystis* have also been reported in various boreal peatlands (Esson et al., 2016; Putkinen et al., 2018). Notably, a very high abundance (66.3%) of USC γ was detected in the alpine meadow. This taxon is more likely to be detected in upland soils that consume atmospheric CH_4 and have pH values > 6.0 (Knief et al., 2003). FWs, an uncultured phylogenetic cluster, was enriched in the marsh. This group, which has been found in plant harvest wetlands, is believed to play a key role in global CH_4 cycling (Zhu et al., 2007).

4.2. Environmental factors shape soil methanotrophic community similarity and phylogenetic similarity more than geographical distance on a small scale

Soil DOC, TC, and TN were all negatively correlated with methanotroph diversity and soil pH had a positive correlation with methanotroph diversity in all the soil samples. The results suggest that methanotrophs are specifically adapted to particular environments, and substrate availability has been suggested to select for type I and type II methanotrophs and limited their growth (Bodelier et al., 2000; Shiao et al., 2018). Soil C/N, TC, TN and soil moisture were the major drivers of distinct community structures, which is supported by findings of a previous study (Kumaresan et al., 2009). In addition, we observed that environmental factors play a critical role in shaping soil methanotrophic community structure and phylogenetic similarity, and geographical distance plays a relatively minor role in determining methanotrophic communities (Table 1; Table S12). This may be because the three habitats are located at a relatively small spatial scale. For comparison, sampling across larger distances should be considered in future studies.

4.3. Stochastic processes dominated in methanotrophic community assembly

Stochastic processes can greatly influence microbial community composition when environmental heterogeneity or species sorting is low (Woodcock et al., 2013; Bennke et al., 2016). In our results, the β NTI values were between -2 and $+2$, which indicated that stochastic processes were shaping methanotrophic community assembly (Fig. 4A and B), and was consistent with the findings of a study conducted in QTP that demonstrated that stochasticity may overwhelm deterministic processes at scales of less than 130 km (Shi et al., 2018). Nevertheless, there were some deterministic processes in the alpine meadow (with a

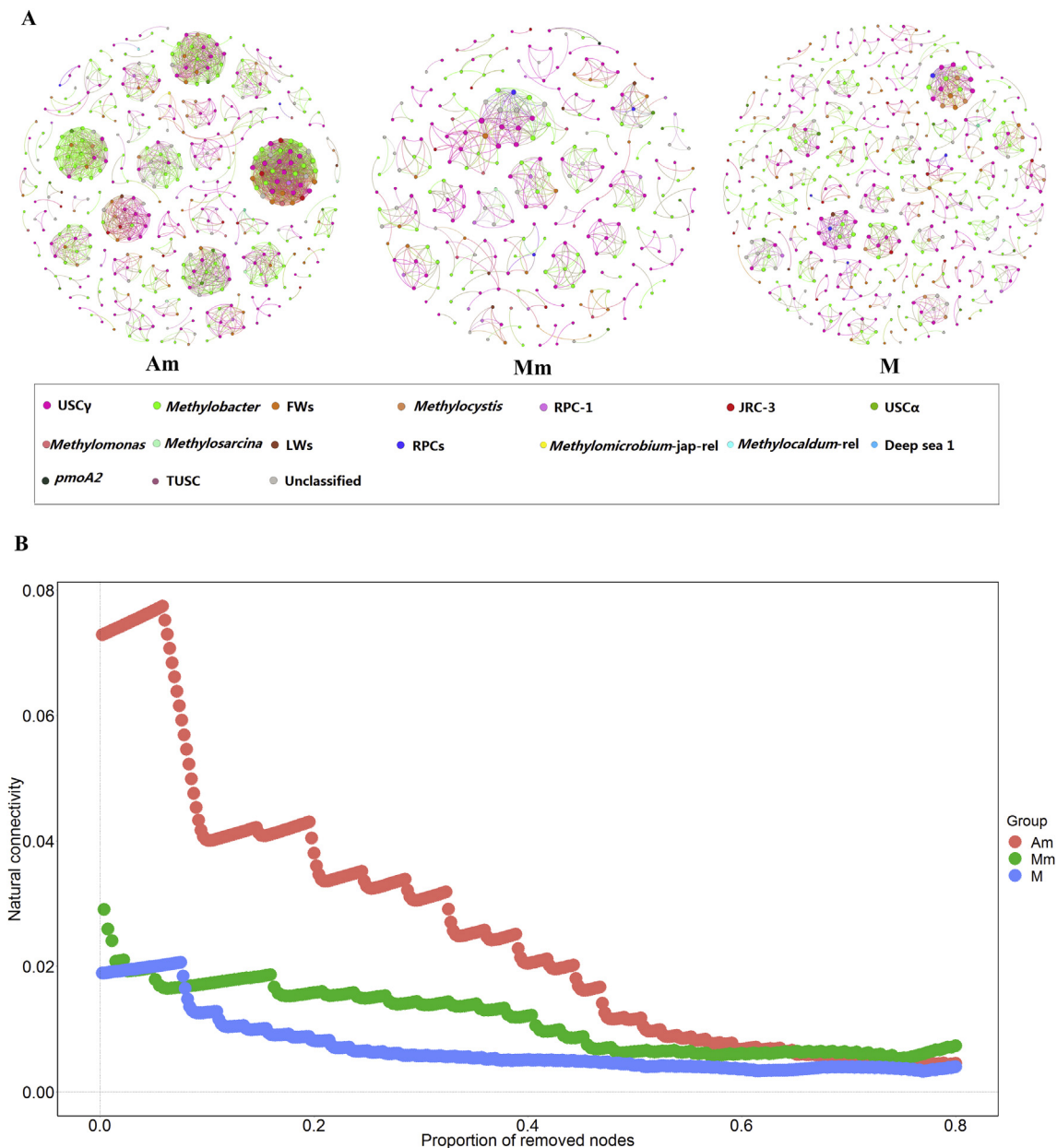


Fig. 5. The co-occurrence network interactions of soil methanotrophic community at the genus level. The size of each node is proportional to the number of connections (i.e., degree). OTUs colored by genus (A); and network stability (B). Am: Alpine meadow; Mm: Marsh meadow; M: Marsh.

relative contribution of 32.4%) and in the marsh meadow (with a relative contribution of 32.0%) compared with the marsh (Table S14). Although the proportion of stochasticity stayed consistent along the water gradients, the result may indicate that methanotroph community composition varied among habitats, which may select for specialist taxa (e.g. atmospheric CH_4 oxidation by $\text{USC}\gamma$). However, our research was constrained within the context of the QTP, and it is important to consider that the relative proportions of methanotrophs in the soil community will shift over space and time – for example with the seasons (Lüke et al., 2014; Smith et al., 2018). More research is required to elucidate the ecological niches of closely related taxa selected along environmental gradients or across space and time. It would be fascinating to know if such stochastic variation produces microscale spatial variation in methane oxidation potential.

4.4. Type I methanotrophs were more responsible for stability of co-occurrence network hubs

Methanotrophic networks exhibited power-law distribution patterns (scale-free networks) in all three habitats, superficially similar to the Internet backbone (Jeong et al., 2000) and metabolic reaction networks (Faloutsos et al., 1999). This means that the power-law distribution networks have many nodes with few links and a few highly connected nodes that are termed as hubs. Networks with these characteristics are considered to be robust towards random node removal but sensitive to the removal of hub nodes (Faust and Raes, 2012). Since the co-occurrence patterns are linked to the interactions of microorganisms in ecosystems, the factors driving the topological features of microbial networks could reflect their effects on microbial interactions (Ma et al., 2016). The environmental drivers of topological features varied among alpine meadow, marsh meadow and marsh networks. Generally, the alpha diversity of soil methanotrophs was more critical than soil

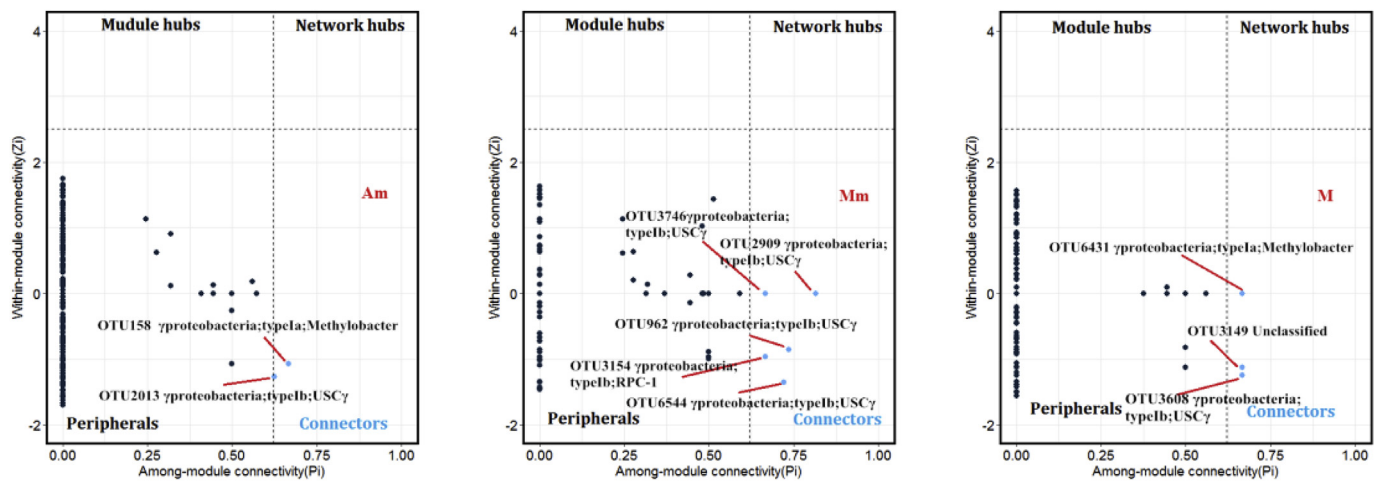


Fig. 6. Z_i - P_i plots showing distribution of OTUs based on their topological roles in methanotrophic networks. Threshold values of Z_i and P_i for categorizing OTUs were 2.5 and 0.62, respectively. Am: Alpine meadow; Mm: Marsh meadow; M: Marsh.

physiochemical variables in influencing subnetwork structures, which could indicate that methanotroph diversity was directly correlated with methanotrophic interactions. We also observed that pH and soil C/N were significantly correlated with subnetwork topological properties in the alpine meadow and the marsh, respectively. The correlations potentially provide some insights for understanding the importance of CH_4 -oxidizing inter-species interactions in influencing element cycling, which could also be influenced by changes in the water tables. From Z_i - P_i space, we suggested that type I methanotrophs (e.g. *Methylobacter*, USC γ , RPC-1) were network connectors and likely to play a key role in community cooperation and have a pronounced impact on the stability of co-occurrence networks. Indeed, several studies indicate that type I methanotrophs are highly active in permafrost wetlands and their roles in low-affinity CH_4 oxidation were varied across soils (Liebner and Wagner, 2007; Yun et al., 2014; Christiansen et al., 2015).

In summary, methanotrophic activity, as observed from our samples tested in laboratory mesocosms, was highest in soils from the marsh and the marsh meadow, as is expected from an environment with higher CH_4 fluxes associated with anoxic conditions in water-saturated environments. Soil methanotrophic assemblages have different niches along a soil water content gradient, resulting in distinct community assemblages in the alpine meadow, the marsh meadow, and the marsh. Although there were different trends in methanotrophic community composition along soil water gradients, the community assembly of methanotrophs appeared to be dominated more by stochastic processes than by deterministic processes, and marshes had a large proportion of stochasticity (Homogenizing dispersal). For the network structure patterns, type I methanotrophs seems to be more responsible for the inter-species cooperation. Collectively, the results suggested that the soil methanotrophic community assemblage processes varied along water table gradients in QTP. Future research should focus on the effect of alternating dry and wet conditions on methanotrophic community to examine community responses and the resilience of CH_4 oxidation.

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Appendix A. Supplementary data

Supplementary data to this article can be found online at <https://doi.org/10.1016/j.soilbio.2019.02.007>.

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